

The Origin of Life—Did It Occur at High Temperatures?

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Received: 10 June 1995 / Accepted: 31 July 1995

Abstract. A high-temperature origin of life has been proposed, largely for the reason that the hyperthermophiles are claimed to be the last common ancestor of modern organisms. Even if they are the oldest extant organisms, which is in dispute, their existence can say nothing about the temperatures of the origin of life, the RNA world, and organisms preceding the hyperthermophiles. There is no geological evidence for the physical setting of the origin of life because there are no unmetamorphosed rocks from that period. Prebiotic chemistry points to a low-temperature origin because most biochemicals decompose rather rapidly at temperatures of 100°C (e.g., half-lives are 73 min for ribose, 21 days for cytosine, and 204 days for adenine). Hyperthermophiles may appear at the base of some phylogenetic trees because they outcompeted the mesophiles when they adapted to lower temperatures, possibly due to enhanced production of heat-shock proteins.

Key words: Origin of life — Hyperthermophiles — Heat shock proteins — Submarine hydrothermal vents — Pre-RNA world — RNA world — Primitive atmosphere

Introduction

Although considerable efforts have been made to understand the emergence of the first living systems, we still do not know when and how life originated. Since it is

sometimes possible to correlate major evolutionary changes with environmental conditions, several attempts have been made to infer the conditions in which life arose by studying the oldest known organisms. As summarized by Pace (1991), molecular evolution analysis has suggested that the oldest recognizable prokaryotes are the anaerobic sulfur-reducing chemosynthetic hyperthermophiles, i.e., organisms that grow optimally at 90°C and above. This observation has been interpreted to imply that the first living systems may have originated in a high-temperature environment, such as those found today in deep-sea hydrothermal vents (Holm 1992) although terrestrial hot springs would be equally consistent. This is a simple extrapolation of the growth temperature of extant hyperthermophiles to the origin of life (Fig. 1, dotted line tree). There is no more justification for this extrapolation than for a mesophilic origin (Fig. 1, solid line tree) or an even higher temperature origin (not shown).

How Old Are the Hyperthermophiles?

A thermophilic origin of life is not a new idea. It was first suggested by Harvey (1924) that the first forms of life were heterotrophic thermophiles that had originated in hot springs. The antiquity of hyperthermophiles is now widely accepted not only for archaebacteria but also for the less well-known eubacterial extremophiles (Stetter 1994). However, alternative opinions have developed since Pace's article in 1991. For example, the archaea may not be the oldest lineage (Doolittle 1995), and there are some indications for separation of the three domains as late as 2×10^9 years ago (Doolittle et al., 1996). In

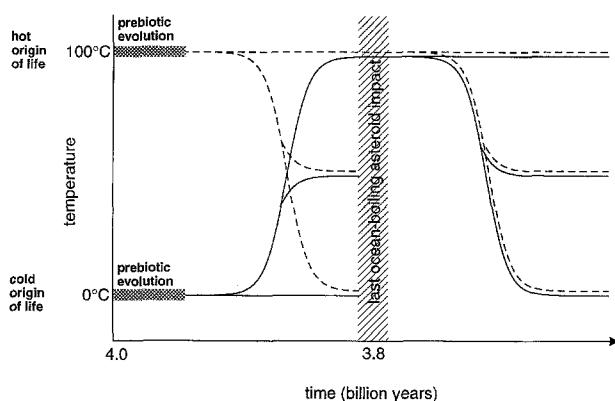


Fig. 1. Some alternative temperature regimes for the origin and early evolution of life. The dotted line shows a hot origin of life followed by adaptation to several lower temperatures. Only the hyperthermophiles would survive the last asteroid impact which boiled the ocean. The solid line shows a low temperature origin of life followed by adaptation to higher temperatures, with survival of only the secondarily adapted hyperthermophiles after the asteroid collision. Not shown are earlier collisions that may have frustrated the origin of life or its survival.

addition, some hyperthermophile sequences are displaced from their basal position if molecular markers other than elongation factors or ATPase subunits are employed (Forterre et al. 1993; Klenk et al. 1994), i.e., other molecular trees also open up the possibility that the last common ancestor of all living beings was not a thermophilic prokaryote. It could be argued that the discovery of frigid archaea living in Antarctic waters (DeLong et al. 1994) supports a low-temperature origin of life, but the same considerations of Fig. 1 apply to these organisms. Thus, the base of the tree at 0° would correspond to the growth conditions of low-temperature archaea (~5°C, instead of 100 to 110°C).

Hyperthermophiles May Be Ancient, But They Are Hardly Primitive

It is important to distinguish between ancient and primitive. Hyperthermophiles may be cladistically ancient, but they are hardly primitive relative to the first living organisms. In fact, they seem to be no more primitive in their metabolic apparatus than mesophiles (Adams 1993). For example, biosynthesis of arginine in *Thermotoga maritima* appears to be essentially the same as in all known organisms (Van de Casteele et al. 1990), and tryptophan biosynthetic genes in the thermoacidophile *Sulfolobus solfataricus* are homologous to their mesophilic eubacterial and eukaryotic counterparts (Tutino et al. 1993).

Primitive living systems, according to some current opinion, would initially refer to the pre-RNA world, in which life was first based on polymers using backbones other than ribose-phosphate and possibly bases different from AUGC. This was followed by a stage in which life was based on RNA as both the genetic material and as

catalysts. The RNA world was followed by a DNA/protein world with rather limited biosynthetic capabilities, and in the subsequent stages of biological evolution, the basic characteristics of metabolic pathways were established. In view of this enormous metabolic development, a constant-temperature extrapolation is hardly justified. These considerations apply if the alternative hypothesis is correct that mesophiles are the most ancient, as suggested by some phylogenetic trees. In this case, constant-temperature extrapolation points to a low-temperature origin of life, but a high-temperature regime or a colder one are equally justified (Fig. 1).

The antiquity of hyperthermophiles fits in with the plausible hypothesis of impact frustration of the origin of life (Maher and Stevenson 1988; Sleep et al. 1989), for which, however, there is no geological evidence. If the last large asteroid to hit the Earth was 400 km in diameter or larger it would have converted the entire ocean to steam. This would have killed off most organisms but the hyperthermophiles would have been selected for, therefore explaining their basal position in some phylogenetic trees. Such extreme thermophiles are sometimes said to be submarine vent organisms (Gogarten-Boekels et al. 1995), but any hyperthermophile would have survived.

It should be kept in mind that even if the whole ocean boils, there may be areas where the temperature is not raised, or where the organisms are protected, e.g., several kilometers deep in marine and continental sediments. Since a boiling ocean would cool in ~1,000 years, the high temperatures would not reach the organisms this deep. Prokaryotic biotas are found today at depths of 0.5 km or more in the oceanic sediments (Parkes et al. 1994), and if such communities were present at the last impact, it is reasonable to think that they would have survived.

What Was the Physical Setting of the Origin of Life?

Many strong statements have been made about the primitive Earth, but there is no direct geological evidence for any of these hypotheses, since there are no rocks older than 3.8×10^9 years. Calculations of atmospheric and planetary formation models, however meritorious, do not constitute evidence.

The temperature of the primitive Earth during the period of the origin of life is unknown. The entire planet is generally thought, without direct evidence, to have remained molten for several hundred million years after its formation 4.6×10^9 years ago (Wetherill 1990). The oldest sedimentary rocks in the Greenland Isua formation have been heated to 500°C, so the evidence on the conditions at that time has been largely destroyed. The sediments in the Australian Warrawoona formation 3.5×10^9 years old contain very convincing cyanobacteria-like microfossils (Schopf 1993).

Thus, life is thought to have originated some time

between 4.0 and 3.5 billion years ago, but there is no direct evidence for the temperature or atmospheric conditions. Some atmospheric models incorporate high partial pressures of CO₂ to raise the temperature of the Earth by a greenhouse effect and thus prevent the complete freezing of the oceans (Kasting 1993). However, a frozen Earth has some advantages for prebiotic chemistry (Bada et al. 1994). But again, there is no direct evidence either way. In addition, processes relevant to the origin of life may have taken place in environments different from the terrestrial average, e.g., hot springs, eutectic sea water, or drying lagoons.

The composition of the atmosphere poses a similar problem: there is no geological evidence as to whether the Earth's atmosphere was reducing or nonreducing, although it is generally accepted that O₂ was absent. It is beyond the scope of this review to explore this question, except to comment that atmospheric chemists mostly favor high CO₂ + N₂, whereas prebiotic chemists mostly favor more reducing conditions (CH₄ + N₂ or CO₂ + H₂ + N₂). Reducing conditions are required for the synthesis of amino acids, purines, pyrimidines, and sugars, and such syntheses are very efficient (Stribling and Miller 1987). The robustness of this type of chemistry is supported by the occurrence of most of these biochemical compounds in the Murchison meteorite, which comes from an asteroid. The meteorite analysis results make it plausible, but do not prove, that such syntheses also occurred on the primitive Earth. Based on what is known about prebiotic chemistry, if the Earth was not reducing, then the organic compounds would have to be brought to it by dust particles, comets, and meteorites (Anders 1989; Chyba et al. 1990). The amounts that can be brought in this way and survive passage through the atmosphere are quite small, and may not have been sufficient for the origin of life.

High Temperatures Give Higher Reaction Rates, But There Is a Price to Pay

The one advantage of high temperatures is that the chemical reactions could go faster and the primitive enzymes could have been less efficient (Harvey 1924), but the price paid is loss of organic compounds by decomposition and diminished stability of the genetic material. The problem with monomers is bad enough, but it is worse with polymers, e.g., RNA and DNA (Lindahl 1993), whose stability in the absence of efficient repair enzymes is too low to maintain genetic integrity in hyperthermophiles. RNA and DNA are clearly too unstable to exist in a hot prebiotic environment. The existence of an RNA world with ribose appears to be incompatible with the idea of a hot origin of life. The stability of ribose and other sugars is the worst problem, but pyrimidines and purines and some amino acids are nearly as bad. The half-life of ribose at 100°C, and pH 7 is only 73 min, and

other sugars have comparable half-lives (Larralde et al. 1995). The half-life for deamination of cytosine at 100°C is 21 days (Garrett and Tsau 1972; Shapiro and Klein 1966), and 204 days at 100°C for adenine (Frick et al. 1987; Shapiro 1995). Some amino acids are stable, e.g., alanine with a half-life for decarboxylation of approximately 19,000 years at 100°C, but serine decarboxylates to ethanolamine with a half-life of 320 days (Vallentyne 1964), with dealdolization and dehydration as additional decomposition routes (Bada et al. 1995). Similar considerations show that the growth of organisms at 250°C or 350°C and the origin of life at such temperatures (Corliss et al. 1981) are very unlikely (White 1984; Miller and Bada 1988). It is clear that if the origin of life took place at 100°C or higher temperatures, then the organic compounds involved must have been used immediately after their prebiotic synthesis. An alternative is to assume an autotrophic origin of life, i.e., the first organisms made all their cell material from CO₂, N₂, and H₂O. Proposals of an autotrophic origin of life are periodically resurrected, but they are made without supporting experiments.

Hyperthermophily May Be a Secondary Adaptation?

An analysis of the rooted tree provided by Stetter (1994) suggests to us a polyphyletic origin of mesophiles from hyperthermophiles, i.e., an independent, parallel abandonment of hyperthermophilic traits has taken place in widely separated branches of the universal tree. This could suggest that relatively few genetic changes may be required to abandon a thermophilic lifestyle. However, other rootings of universal trees permit the evolution of mesophilic branches into hyperthermophilic ones (Rivera and Lake 1992; Klenk et al. 1994; Forterre et al. 1993).

A possibility that has not been given enough attention is that hyperthermophiles are now at the base of some trees simply because they outcompeted older mesophiles when they adapted to lower temperatures, rather than being the sole survivors of an impact event. Some of the molecular features that are adaptations to hot environments could have enhanced the survival chances of hyperthermophiles and their immediate descendants under less-extreme temperature conditions. An example could be the heat-shock response, whose universal distribution can be interpreted as a remnant of the hyperthermophilic ancestors of extant life. Heat-shock proteins are not only involved in thermotolerance, but also in protection against other stress-inducing agents and environmental insults, including starvation conditions, UV-irradiation, DNA-damaging agents, alcohol, amino acid analogues, etc. (Watson 1990). Accordingly, it is possible to envision that heat-shock genes evolved in ancient hyperthermophiles, preadapting them to other stress-inducing

conditions at low temperatures, allowing them to successfully outcompete mesophiles. While this scheme may not necessarily be correct, it suggests that additional explanations can be advanced to explain the phylogenetic distribution of hyperthermophiles.

We have addressed the possibility that hyperthermophiles are not the oldest organisms. Even if they are, a straight-line temperature extrapolation back in time to the origin of life is not warranted. Prebiotic chemistry points toward a low-temperature regime for the emergence of living systems. If this conclusion is valid, it merits a search for mesophiles older than hyperthermophiles.

Acknowledgments. We thank Jason P. Dworkin and Jeffrey Bada for helpful comments. Support was provided by the NSCORT (NASA Specialized Center for Research and Training) in Exobiology at the University of California, San Diego.

References

- Adams MWW (1993) Enzymes and proteins from organisms that grow near and above 100°C. *Annu Rev Microbiol* 47:627–658
- Anders E (1989) Pre-biotic organic matter from comets and asteroids. *Nature* 342:255–257
- Bada JL, Bigham C, Miller SL (1994) Impact melting of frozen oceans on the early Earth: implications for the origin of life. *Proc Natl Acad Sci USA* 91:1248–1250
- Bada JL, Miller SL, Zhao M (1995) The stability of amino acids at submarine hydrothermal vent temperatures. *Orig Life Evol Biosph* 25:111–118
- Chyba CF, Thomas PJ, Brookshaw L, Sagan C (1990) Cometary delivery of organic molecules to the early Earth. *Science* 249:366–373
- Corliss, JB, Baross JA, Hoffman SE (1981) An hypothesis concerning the relationship between submarine hot springs and the origin of life on Earth. *Oceanologica Acta* 4, Suppl 59–69
- DeLong EF, Wu KY, Prézelin BB, Jovine RVM (1994) High abundance of Archaea in Antarctic marine picoplankton. *Nature* 371:695–697
- Doolittle RF (1995) Of Archae and Eo: what's in a name? *Proc Natl Acad Sci USA* 92:2421–2423
- Doolittle RF, Feng DF, Tsang S, Cho G, Little E (1996) Determining the divergence times of the major kingdoms of living organisms with a protein clock. *Science* (submitted)
- Forterre P, Benachenhou-Lahfa N, Confalonieri F, Duguet M, Elie Ch, Labedan B (1993) The nature of the last universal ancestor and the root of the tree of life, still open questions. *Biosystems* 28:15–32
- Frick L, Mac Neela JP, Wolfenden R (1987) Transition state stabilization by deaminases: rates of nonenzymatic hydrolysis of adenosine and cytidine. *Bioorg Chem* 15:100–108
- Garrett ER, Tsau J (1972) Solvolyses of cytosine and cytidine. *J Pharm Sci* 61:1052–1061
- Gogarten-Boekels M, Hilario E, Gogarten JP (1995) The effects of heavy meteoritic bombardment on the early evolution—the emergence of the three domains of life. *Orig Life Evol Biosph* 25:251–264
- Harvey RB (1924) Enzymes of thermal algae. *Science* 60:481–482
- Holm NG ed (1992) Marine hydrothermal systems and the origin of life. Kluwer, Academic, Dordrecht. 242 pp. Also a special issue of *Origins of Life and Evolution of the Biosphere* 22:1–241
- Kasting JF (1993) Earth's earliest atmosphere. *Science* 259:920–926
- Klenk HP, Palm P, Zillig W (1994) DNA-dependent RNA polymerases as phylogenetic marker molecules. *Syst Appl Microbiol* 16:638–647
- Larralde R, Robertson MP, Miller SL (1995) Rates of decomposition of ribose and other sugars: implications for chemical evolution. *Proc Natl Acad Sci USA* 92:8158–8160
- Lindahl T (1993) Instability and decay of the primary structure of DNA. *Nature* 362:709–715
- Maher KA, Stevenson DJ (1988) Impact frustration of the origin of life. *Nature* 331:612–614
- Miller SL, Bada JL (1988) Submarine hot springs and the origin of life. *Nature* 334:609–611
- Pace NR (1991) Origin of life—facing up to the physical setting. *Cell* 65:531–533
- Parkes RJ, Cragg BA, Bale SJ, Getliff JM, Goodman K, Rochelle FA, Fry JC, Weightman AJ, Harvey SM (1994) Deep bacterial biosphere in Pacific Ocean sediments. *Nature* 371:410–413
- Rivera MC, Lake JA (1992) Evidence that eukaryotes and eocyte prokaryotes are immediate relatives. *Science* 257:74–75
- Shapiro R (1995) The prebiotic role of adenine: a critical analysis. *Orig Life Evol Biosph* 25:83–98
- Shapiro R, Klein RS (1966) The deamination of cytidine and cytosine by acidic buffer solutions. Mutagenic implications. *Biochemistry* 5:2358–2362
- Schopf JW (1993) Microfossils of the early Archean Apex chert: new evidence of the antiquity of life. *Science* 260:640–646
- Sleep NH, Zahnle KJ, Kasting JF, Morowitz HJ (1989) Annihilation of ecosystems by large asteroid impacts on the early Earth. *Nature* 342:139–142
- Stetter KO (1994) The lesson of archaeobacteria. In: Bengtson S. (ed) *Early life on earth: Nobel symposium No. 84*. Columbia University Press, New York, pp 114–122
- Stribling R, Miller SL (1987) Energy yields for hydrogen cyanide and formaldehyde synthesis: the HCN and amino acid concentrations in the primitive oceans. *Orig Life Evol Biosph* 17:261–273
- Tutino ML, Scarano G, Marino G, Sanna G, Cubellis MV (1993) Tryptophan biosynthesis genes *trpEGC* in the thermoacidophilic archaeobacterium *Sulfolobus solfataricus*. *J Bacteriol* 175:299–302
- Vallentyne JR (1964) Biogeochemistry of organic matter II: thermal reaction kinetics and transformation products of amino compounds. *Geochim Cosmochim Acta* 28:157–188
- Van de Castele M, Demarez M, Legrain C, Glansdorff N, Pierard A (1990) Pathways of arginine biosynthesis in extreme thermophilic archaeobacteria and eubacteria. *J Gen Microbiol* 136:1177–1183
- Watson K (1990) Microbial stress proteins. *Adv Microbiol Physiol* 31:183–223
- Wetherill GW (1990) Formation of the Earth. *Annu Rev Earth Planet Sci* 18:205–256
- White RH (1984) Hydrolytic stability of biomolecules at high temperatures and its implication for life at 250°C. *Nature* 310:430–432